

KECK GEOLOGY CONSORTIUM

21ST KECK RESEARCH SYMPOSIUM IN GEOLOGY SHORT CONTRIBUTIONS

April 2008

Dr Andrew P. de Wet, Editor
Keck Director
Franklin & Marshall College

Keck Geology Consortium
Franklin & Marshall College
PO Box 3003, Lancaster Pa, 17603

Dr Amy Rhodes,
Symposium Organizer
Smith College

Keck Geology Consortium Member Institutions:

Amherst College Beloit College Carleton College Colgate University The College of Wooster The Colorado College
Franklin and Marshall College Macalester College Mt. Holyoke College Oberlin College Pomona College Smith College Trinity University
Union College Washington and Lee University Wesleyan University Whitman College Williams College

2007-2008 PROJECTS:

Tectonic and Climatic Forcing of the Swiss Alps

John Garver (Union College), Mark Brandon (Yale University), Alison Anders (University of Illinois),
Jeff Rahl (Washington and Lee University), Devin McPhillips (Yale University)
Students: William Barnhart, Kat Compton, Rosalba Queirolo, Lindsay Rathnow,
Scott Reynhout, Libby Ritz, Jessica Stanley, Michael Werner, Elizabeth Wong

Geologic Controls on Viticulture in the Walla Walla Valley, Washington

Kevin Pogue (Whitman College) and Chris Oze (Bryn Mawr College)
Students: Ruth Indrick, Karl Lang, Season Martin, Anna Mazzariello, John Nowinski, Anna Weber

The Árnes central volcano, Northwestern Iceland

Brennan Jordan (University of South Dakota), Bob Wiebe (Franklin & Marshall College), Paul Olin (Washington State U.)
Students: Michael Bernstein, Elizabeth Drewes, Kamilla Fella, Daniel Hadley, Caitlyn Perlman, Lynne Stewart

Origin of big garnets in amphibolites during high-grade metamorphism, Adirondacks, NY

Kurt Hollocher (Union College)
Students: Denny Alden, Erica Emerson, Kathryn Stack

Carbonate Depositional Systems of St. Croix, US Virgin Islands

Dennis Hubbard and Karla Parsons-Hubbard (Oberlin College), Karl Wirth (Macalester College)
Students: Monica Arienzo, Ashley Burkett, Alexander Burpee, Sarah Chamlee, Timmons Erickson
Andrew Estep, Dana Fisco, Matthew Klinman, Caitlin Tems, Selina Tirtajana

Sedimentary Environments and Paleoecology of Proterozoic and Cambrian "Avalonian" Strata in the United States

Mark McMenamin (Mount Holyoke College) and Jack Beuthin (U of Pittsburgh, Johnstown)
Students: Evan Anderson, Anna Lavarreda, Ken O'Donnell, Walter Persons, Jessica Williams

Development and Analysis of Millennial-Scale Tree Ring Records from Glacier Bay National Park and Preserve, Alaska (Glacier Bay)

Greg Wiles (The College of Wooster)
Students: Erica Erlanger, Alex Trutko, Adam Plourde

The Biogeochemistry and Environmental History of Bioluminescent Bays, Vieques, Puerto Rico

Tim Ku (Wesleyan University) Suzanne O'Connell (Wesleyan University), Anna Martini (Amherst College)
Students: Erin Algeo, Jennifer Bourdeau, Justin Clark, Margaret Selzer, Ulyanna Sorokopoud, Sarah Tracy

Funding provided by:

Keck Geology Consortium Member Institutions and NSF (NSF-REU: 0648782)

**Keck Geology Consortium: Projects 2007-2008
Short Contributions – Avalonian**

**FINE CLASTICS OF THE BOSTON BAY GROUP: NEW DATA AND
INTERPRETATIONS CONCERNING DEPOSITIONAL PROCESSES AND
ENVIRONMENTS: p 209-212**

Project faculty:

MARK A. S. McMENAMIN: Mount Holyoke College

JOHN D. BEUTHIN: University of Pittsburgh, Johnstown

**A PETROGRAPHIC AND SEM-EDS ANALYSIS OF *ASPIDELLA*-BEARING
SILTSTONES AND SLATES OF THE CAMBRIDGE ARGILLITE, BOSTON
BAY GROUP, MASSACHUSETTS: p213-218**

EVAN PELZNER ANDERSON: University of Mary Washington

Research Advisor: Jodie L. Hayob

**LAMINATED CAMBRIDGE ARGILLITE OF THE BOSTON BAY GROUP AT
HEWITT'S COVE, MASSACHUSETTS: p219-222**

ANNA E. LAVARREDA: Smith College

Research Advisor: Bosiljka Glumac

**ENVIRONMENTAL ANALYSIS OF THE NEOPROTEROZOIC CAMBRIDGE
ARGILLITE, BOSTON BASIN, MASSACHUSETTS: p223-228**

KENNETH O'DONNELL: Beloit College

Research Advisor: Carl Mendelson

**A FIELD AND LABORATORY STUDY OF THE EDIACARAN FOSSILS OF
HEWITT'S COVE: EVIDENCE OF TECTONIC DEFORMATION AND
CONSIDERATION OF PALEOBIOLOGY: p229-233**

W. SCOTT PERSONS, IV: Macalester College

Research Advisor: Raymond Rogers

**LAMINITES AND DROPSTONES IN THE CAMBRIDGE ARGILLITE
(EDIACARAN), HEWITT'S COVE, HINGHAM, MASSACHUSETTS: p234-237**

JESSICA A. J. WILLIAMS: Southern Utah University

Research Advisor: C. Frederick Lohrengel II

Funding provided by: Keck Geology Consortium Member Institutions and NSF (NSF-REU: 0648782)

Keck Geology Consortium
Franklin & Marshall College
PO Box 3003, Lancaster Pa, 17603
Keckgeology.org

A FIELD AND LABORATORY STUDY OF THE EDIACARAN FOSSILS OF HEWITT'S COVE: EVIDENCE OF TECTONIC DEFORMATION AND CONSIDERATION OF PALEOBIOLOGY

W. SCOTT PERSONS, IV: Macalester College
Academic Advisor: Raymond Rogers

INTRODUCTION

Ediacaran fossils have baffled scientists since their original discovery in 1868. As the oldest unambiguously multicellular organisms with tissues, Ediacaran organisms (or simply Ediacarans) represent a key stage in the evolution of life, and understanding their fossil record is essential to understanding the status of life prior to the great radiation of forms seen in the Early Cambrian. However, the bizarre shapes and structures of many Ediacarans have made placing them in phylogenetic context with modern taxa and determining their ecological roles difficult. In this study, the Ediacaran fossils of a roughly 575 Ma argillite bed of the Cambridge Formation, exposed at Hewitt's Cove, Massachusetts (see Fig. 1), were described and considered in relation to current theories on Ediacaran paleobiology.



Figure 1. The inclined surfaces of the argillite bed at Hewitt's Cove, 2007.

THE EDIACARAN *ASPIDELLA*

Despite careful combing of the field site and microscopic examination of numerous thin sections, no evidence of any fossil genus, other than the Ediacaran *Aspidella*, was found. *Aspidella* are mysterious organisms, even by Ediacaran standards. Specimens consist of a fossil structure that superficially resembles a small crater, generally no more than a few centimeters in diameter and a few millimeters in height (see Fig. 2). The name '*Aspidella*' was first used by Elkanah Billings in 1872 to describe ring-shaped fossils from Newfoundland (Billings, 1872). The type specimen of *Aspidella* has been lost (Gehling et al., 2000), leaving the genus vulnerable to nomenclatural uncertainty. *Aspidella* has since be-

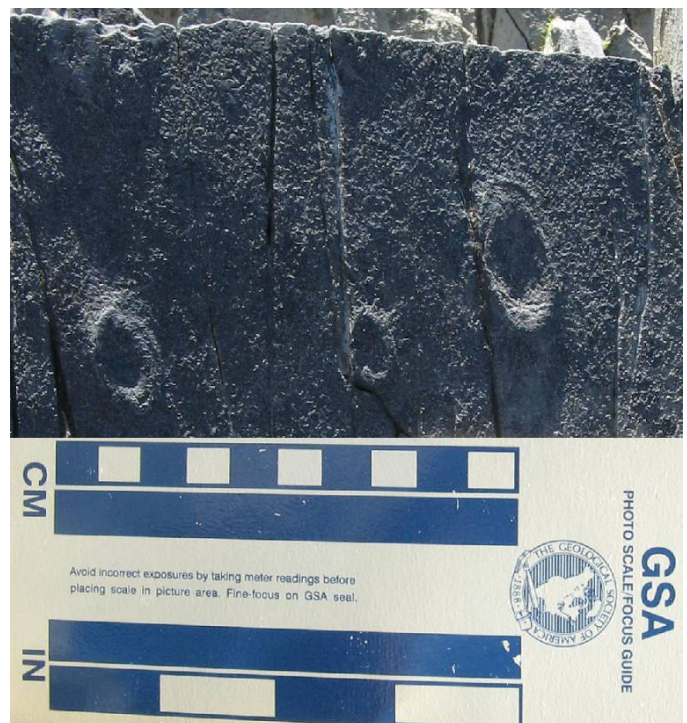


Figure 2. *Aspidella* in outcrop at Hewitt's Cove.

come somewhat of a junkyard taxon, in which many varieties of ring-shaped fossils with otherwise indeterminate morphologies have been lumped (Gehling et al., (2000).

The original description of *Aspidella* suggests the type specimen actually had multiple ridges radiating from the center of the ring, like the spokes of a bicycle wheel (Gehling et al., 2000). If this is true, the name *Aspidella* is actually more morphologically specific than commonly perceived. However, for the purposes of this study, the name *Aspidella* will be used in the standard way (as a catch-all genus for ring-shaped Ediacarans), with the understanding that the genus is in need of revision.

The nature of *Aspidella* is the subject of much discussion and debate. The general consensus is that they were relatively sessile organisms, similar in appearance, but with perhaps no evolutionary relation, to modern members of the order Pennatulacea (the sea pens). They, and many other Ediacaran taxa, are believed to have each had a soft stalk (the rachis) with a frond-like appendage at one end and a stiff base (the peduncle) at the other (terminology from Laflamme and Narbonne, 2008). This base anchored the *Aspidella* to the sea floor. Only the stiff peduncles of most *Aspidella* are preserved as fossils, but comparison of the bases to those of other fossil organisms gives credence to the notion that stalks and fronds were present in life. In rare instances, *Aspidella* specimens from the Fermeuse Formation have been found with what appear to be impressions of stalks attached to the more clearly preserved peduncles (Gehling et al., 2000).

That no fossils other than those of *Aspidella* were discovered does not imply that only one variety of Ediacaran was present in Precambrian Hewitt's Cove. Given that only a portion of each *Aspidella* was ever found (never any more than the organism's holdfast structure), it would be expected that any Ediacarans with bodies softer than the holdfasts of *Aspidella* might not be preserved. This effectively rules out nearly all other Ediacaran forms. So, while Hewitt's Cove is a location dominated by a single fossil organism, it may well have been an

environment of far richer biodiversity. Nevertheless, *Aspidella* were the only Ediacarans found during the 2007 field season, and it is solely on them that interpretations of the location's paleoecology must be based.

THE POSSIBILITY OF TECTONIC STRETCHING

Hewitt's Cove's *Aspidella* fossils all have an elliptical shape, and this is an important characteristic to understand. It distinguishes them from the *Aspidella* of most other locations, including some that are temporally congruent, such as those of the Mistaken Point assemblage (Canfield et al., 2007). Because the long axes of the elliptical fossils all have similar orientations, it has been speculated that these *Aspidella* were shallow-water dwellers that evolved a streamlined shape and aligned themselves with currents or light sources for photosynthesis (McMenamin, personal communication, 2007). However, the planar, normally graded, and fine-grained laminations of the argillites, combined with their observed lack of traction features and storm indicators (e.g., hummocy and swaley laminae), suggest they formed below the storm wave-base, where it is unlikely strong currents were present.

Measurements of the orientation of the *Aspidella* show exceedingly uniform orientations throughout the exposure, not just on a single bedding surface. This raises the possibility of another explanation for both the unusual shape of the *Aspidella* and their shared long-axis orientation: tectonic deformation. Ancient marine fossils are commonly known to become stretched and distorted in a particular direction by the stress/strain of tectonic events (Webster and Hughes, 1999; Underwood, 1992; Cooper, 1990). Assuming that the *Aspidella* bases were actually circular when first fossilized, a tectonic event could have stretched the fossils into elliptical shapes and would be expected to stretch the great majority of fossils throughout the local exposure in the same direction and to the same degree. The *Aspidella* fossils of the Conception and St. Johns' Group of the Avalon Peninsula are also elliptical, and research has indicated that tectonic deformation is responsible

for the elliptical shapes in these instances (Gehling et al., 2000).

SIZE DISTRIBUTION OF THE HEWITT'S COVE POPULATION

Given the limited fossil preservation and the inferred tectonic deformation, it is impossible to determine, based on morphology alone, whether all *Aspidella* fossils observed at Hewitt's Cove belong to the same species or the same subspecies (or the same genus, given the suspect nature of the *Aspidella* taxon). However, the holdfast fossils all have identical morphology and give no reason to assume the presence of different species. Figure 3 shows the relative abundance of *Aspidella* fossils within their measured range of total elliptical surface area. The graph shows a unimodal distribution. Were the population composed of more than one *Aspidella* species, a bimodal distribution in size might be expected. Whether or not the population is composed of a single species remains unproven, but given the unimodal distribution and the identical (though poorly preserved) structure of the holdfast fossils, it is most parsimonious to conclude that the argillites of Hewitt's Cove record only one species of *Aspidella*.

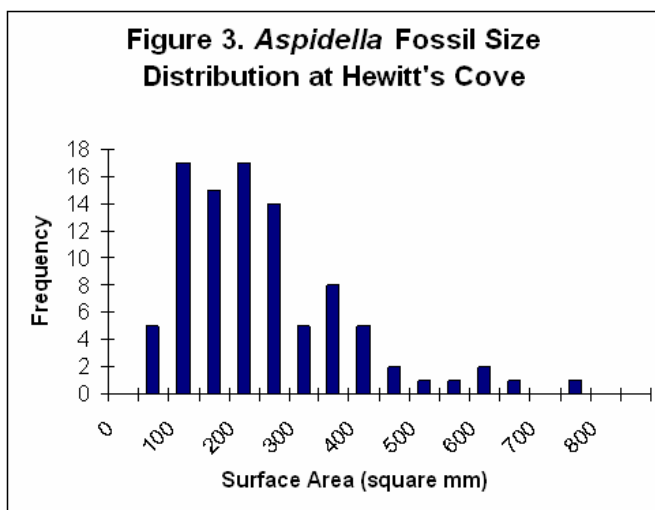


Figure 3. Size distribution of Hewitt's Cove *Aspidella* population (n=95).

As has been previously observed (Bailey and Ross,

1993), small and large *Aspidella* appear side by side on the same surfaces throughout the Hewitt's Cove exposure, and *Aspidella* of any given size appear equally common at all elevations. This suggests that variations in taphonomic sorting were minimal throughout the depositional history of the argillites. However, this does not mean that the *Aspidella* population as a whole was not subject to a single consistent sorting bias or to multiple consistent sorting biases. The size distribution graph is decidedly right skewed, and taphonomic bias against large individuals may explain this shape. Alternatively, this shape may indicate a population where the majority of individuals failed to reach maximum size.

As always with Ediacarans, eagerness to extrapolate should be tempered. Virtually nothing is known about the life history of *Aspidella*. Still, the simplest interpretation is that *Aspidella* had only one morphological stage in which they grew, reproduced, and died. Following this interpretation, the Hewitt's Cove size distribution graph may depict a population in which individuals grew larger over time and were characterized with greater mortality early in life than later.

On the whole, the graph contradicts what would be predicted from theories that propose *Aspidella* to have reproduced by division (Grazhdankin and Gerdes, 2007). If *Aspidella* reproduced simply by splitting themselves in two, the largest individuals should be roughly twice the size of the smallest (assuming perfect one-splits-into-two divisions and that the *Aspidella* consistently split at the first opportunity). Instead, the largest individuals are nearly fifteen times as large as the smallest individuals. If *Aspidella* did reproduce by division, their divisions must have either been unequal or there must have been great variability in the intervening pre-division periods of growth.

EVIDENCE FOR MOVEMENT

Although the argillite beds of the Cambridge Formation record only a limited portion of the *Aspidella* structure, the beds are ideal for answering questions related to *Aspidella* locomotion, because the fossils

are embedded in laminated rock. Were the laminations not present, it would be impossible to detect areas of the argillite that had been disturbed; however, with them, movement of the *Aspidella* (especially vertical movement) should be seen as regions where the laminations are blurred and/or disrupted. All examined thin sections showed no signs of lamination disruption (see Fig. 4). This suggests *Aspidella* had a stationary lifestyle, or were only mobile above the sediment-water interface.

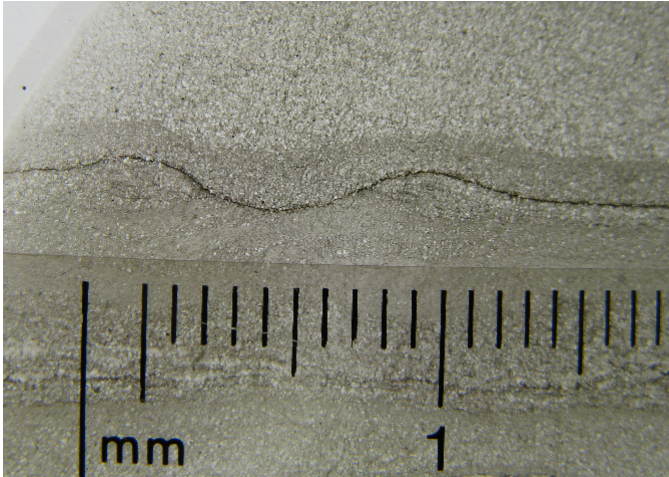


Figure 4. *Aspidella* in thin section, Hewitt's Cove Section.

The lack of lamination disruption above the *Aspidella* fossils indicates that the organisms anchored themselves to the surface of the sea floor or rested upon it. Had they buried their holdfasts any distance below the top layer of sediment, lamination disturbance above the *Aspidella* fossils would be expected. This observation contrasts with models that relate *Aspidella* to modern sea pens – many of which do actively bury their bases (Kastendiek, 1976). The lack of horizontal disturbances around the *Aspidella* indicates that the organisms were also incapable of, or at least seldom engaged in, lateral movement. Finally, the lack of any disturbance around all observed *Aspidella* implies that the organisms were buried without a struggle. There is no indication here -- as there is in some instances of turbidite-buried trilobites (Speyer and Brett, 1986) -- that the *Aspidella* were buried suddenly and attempted to escape. Modern sea pens are known to have the ability to extricate themselves after being buried (Kastend-

iek, 1976). The *Aspidella* were either too immobile to fight through the sediment load or were already dead at the time of burial.

REFERENCES

- Bailey, R. H. and M. E. Ross, 1993. Geology of East Point, Nahant, Massachusetts: field trip guidebook for the Northeastern United States: 1993. Boston GSA, University of Massachusetts, Amherst, pages Y1-Y24.
- Billings, E., 1872. Fossils in Huronian rocks. Canadian Naturalist and Quarterly Journal of Science, vol. 6, page 478.
- Canfield, D. E., S. W. Poulton, and G. M. Narbonne, 2007. Late-Neoproterozoic deep-ocean oxygenation and the rise of animal life. Science, vol. 315, pages 92-95.
- Cooper, R. A. 1990. Interpretation of tectonically deformed fossils. New Zealand Journal of Geology and Geophysics, vol. 33, pages 321-323.
- Gehling, J.G., G. M. Narbonne, and M. M. Anderson, 2000. The first named Ediacaran body fossil, "*Aspidella terranovica*". Palaeontology, vol. 43 (3), pages 427-456.
- Grazhdankin D., and G. Gerdes, 2007. Ediacaran microbial colonies. Lethaia, vol. 40, pages 201-210.
- Kastendiek, J. 1976. Behavior of the sea pansy *Renilla kollikeri pfeffer* (Coelenterate: Pennatulacea) and its influence on the distribution and biological interactions of the species. Biol. Bull., vol. 151, pages 518-537.
- Laflamme, M., and G. M. Narbonne. 2008. Ediacaran fronds. Palaeogeography, Palaeoclimatology, Palaeoecology, vol. 18, pages 162-179.
- Speyer, S. E., and C. E., Brett, 1986. Trilobite taphonomy and Middle Devonian taphofacies. Palaios, vol. 1 (3), Theme Issue: Taphonomy: Ecology's

Loss is Sedimentology's Gain, pages 312-327.

Underwood, C. J., 1992. Graptolite preservation and deformation. *Palaios*, vol. 7 (2), pages 178-186.

Webster M., N. C. Hughes, 1999. Compaction-related deformation in Cambrian olenelloid trilobites and its implications for fossil morphometry. *Journal of Paleontology*, vol. 73 (2), pages 355-371.